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Diversity patterns of the vascular plant group Zosterophyllopsida in relation to Devonian paleogeography

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ABSTRACT

The Zosterophyllopsida originated in the Silurian and became prominent vascular components of Early Devonian floras worldwide. An updated dataset of zosterophyllopsids at species level is analyzed to compare the taxic composition of five putative paleophytogeographic units, Laurussia, Siberia, northwestern Gondwana, Kazakhstan and northeastern Gondwana (i.e., Australia, China and the Shan–Thai block). The high level of endemism shown by each unit confirms the phytogeographic differentiation and the occurrence of geographical barriers preventing massive floral exchanges between the corresponding regions for the Late Silurian–Early Devonian time interval. Statistical analyses conducted on the three largest datasets, those corresponding to Laurussia, Siberia and northeastern Gondwana, indicate that the diversity dynamics of the group followed the same pattern in these regions. Almost all taxic diversity measures show that specific diversity was greatest in the middle and late Pragian. Diversity dropped significantly thereafter; however, residual diversity reveals genuine regional patterns. From this, we show that the radiation of the Zosterophyllopsida may have stopped earlier in northeastern Gondwana and Siberia than Laurussia. We propose that the onset of these extinctions resulted from the competitive replacement of the zosterophyllopsids by increasingly diversified lycopsids and basal euphyllophytes whose evolution could have been favored by external factors.

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1. Introduction

The earliest known appearance of the zosterophylls dates back to the Ludfordian stage (Ludlow series, late Silurian) about 425 million years ago, with the earliest macro-remains found in Bathurst Island, Canadian Arctic (Kotyk et al., 2002). With their lateral sporangia arranged with varying density on the stem and a tendency for pseudomonopodial branching and rhizomatous habit (Fig. 1), the zosterophylls appear morphologically more complex than most basal tracheophytes of rhyniophyte/rhyniophytoid affinities (e.g., *Cooksonia*) that preceded them by about 5 million years in the fossil record (Banks, 1968; Edwards et al., 1992; Gensel, 1992; Edwards and Richardson, 2004; Genez and Gerrienne, 2010). The diversity of the zosterophylls in terms of absolute number of genera increased in the Early Devonian. In percentage, they represented the second most diverse group in the Late Silurian and the Lochkovian; they became the dominant group of vascular plants in the Pragian and Emsian (Hao and Xue, 2013).

When Banks (1968, 1975) recognized the Zosterophyllophytina as a subdivision of the Tracheophyta, this group comprised a single order and a single family with six genera, *Zosterophyllum* Penhallow 1892,

Gosslingia Heard 1927, *Crenaticaulis* Edwards ex Banks & Davis 1969, *Sawdonia* Hueber 1971, *Rebuchia* Hueber 1971 and *Bathurstia* Hueber 1972. At this time zosterophylls were already known to have been widely distributed, with macrofossils recorded from many localities in Europe, but also North America, Canadian Arctic, Siberia, Ukraine, South China, Vietnam and Australia (see SI in Cascales-Miñana and Meyer-Berthaud, 2014). Today, more than thirty genera are assigned to the zosterophylls. The most numerous additions have come from South China, Belgium and Great Britain (Gerrienne, 1996a, 1996b; Edwards and Richardson, 2004; Hao and Xue, 2013). Except for Xinjiang in China (see Hueber and Banks, 1979; Cai et al., 1993; Wang and Hao, 1996; Xu et al., 2011) zosterophylls have not been recorded in any new area since Bank's (1975) paper. Outside South China, the Gondwanan record of zosterophylls remains very limited.

With the advent of phylogenetic systematics using large datasets, a number of proposals challenging Bank's conclusions have been made regarding the nature of the zosterophylls and their evolutionary patterns of diversification (Gensel, 1992; Kenrick and Crane, 1997; Hao and Xue, 2013). All include *Sawdonia*, *Gosslingia* and related genera in a zosterophyllopsid clade. The Zosterophyllopsida sensu Kenrick and Crane (1997) is circumscribed to taxa that have circinate growth and a two-rowed arrangement of sporangia; it includes the heterosporous genera *Barinophyton* and *Protobarinophyton*. The zosterophyllopsid

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Fig. 1. Whole plant reconstructions highlighting morphological diversity within the Zosterophylloids. (A) *Zosterophyllum myretonianum*. Upright axes branch sparsely and are naked; sporangia are arranged helically, in loose spikes. Reprinted with permission from Cleal and Thomas (1999, p. 20). (B) *Sawdonia ornata*. B.1: upright axes branch frequently, equally or unequally, and bear spines. B.2: sporangia are arranged in loose spikes. Reprinted with permission from Cleal and Thomas (1999, p. 20). (C) *Oricilla bilinearis*. C.1: upright axes branch moderately, equally or unequally, have a circinate tip, and are naked. Modified by Annette Townsend from Gensel (1982, p. 347). C.2: sporangia are auricular and arranged in 2 rows. Modified by Annette Townsend from Gensel (1982, p. 347). (D) *Serrulacaulis furcatus*. Upright axes branch equally, have a circinate tip, and bear deltoid emergences in two rows; horizontal axes produce rhizoids; sporangia are borne in two rows on one side of upright axis. Modified by Annette Townsend from Hueber and Banks (1979, p. 185).

clade resolved in Hao and Xue's (2013) analyses excludes all heterosporous genera; it comprises taxa with sporangia borne helically and eventually positioned terminally on the stems.

In a recent paper (Cascales-Miñana and Meyer-Berthaud, 2014), we investigated the apparent dynamics of diversification of the Zosterophylloids *sensu* Hao and Xue (2013) using diversity metrics and rates of evolution. We showed that zosterophylloid diversity exhibits a quick increase from the Late Silurian to the end-Pragian, followed by a marked fall ending in the Eifelian. So, during the Ludfordian–Pragian time interval, this pattern of apparent plant diversity changes (i.e., based on data taken at face value; Foote, 2001) is consistent with that predicted by Sepkoski's (1978) diversity-dependent model of diversification in its first phase. However, it does not show the equilibrium phase following the peak of diversity expected by the model in the absence of any major perturbation.

Despite the warnings of Meyen (1987) and Edwards (1990) about the limitations of the exercise, some phytogeographical patterns were recognized for early macrofloral assemblages. For the Late Silurian

Raymond et al. (2006) distinguished four phytogeographical units, north Laurussia (Bathurst Island), Kazakhstan, Australia and a large south Laurussian–northwest Gondwanan unit. Edwards and Berry (1991) proposed a possible five-partite pattern of Early Devonian floral assemblages corresponding to Australia, China, Kazakhstan, Siberia and Laurussia, the three latter ones being consistent with Petrosyan's (1967) results concerning the former Soviet Union. The statistical analysis carried out on plant traits by Raymond (1987) provided convergent results. It clearly distinguished the Australian and Kazakhstan units and, to a lesser degree, China, from a large so-called “equatorial and middle latitude unit” comprising Siberia and Laurussia. Hao and Gensel (1998) recognized the distinctiveness of the floral assemblages from South China and Australia and proposed their grouping in a northeastern Gondwanan phytogeographical unit lying between the paleoequator and 15°S during the Early Devonian.

If paleogeography may have had a role in shaping plant paleodiversity, no phytogeographical analyses have been dedicated to any particular early vascular plant group. Regarding the Zosterophylloids, Raymond et al. (2006), like Edwards and Richardson (2004) previously, recognized a preponderance of the zosterophylloid diversity in north Laurussia by contrast to the rhyniophytes and rhyniophytoids that flourished elsewhere. They proposed that this distinction may have corresponded to a climatic rather than purely geographic (i.e., related to paleogeographic barriers for example) control of diversity. In the Late Silurian, whereas zosterophylloids preferentially occupied paleoequatorial regions, rhyniophytes and rhyniophytoids may have been better adapted to the seasonal climates of the subequatorial and temperate paleolatitudes (Edwards and Richardson, 2004; Raymond et al., 2006).

Hao and Xue (2013) noted the large amount of data recently accumulated on early land plant taxa that should stimulate new phytogeographical studies. In this contribution, we provide an updated dataset at species level of the paleogeographical distribution of the Zosterophylloids. We test the relevance of the phytogeographical differentiation hypothesized by the authors cited above (Petrosyan, 1967; Edwards and Berry, 1991; Hao and Gensel, 1998) and we examine to what extent paleogeography may have contributed to the dynamics of the zosterophylloid diversity in the Early Devonian.

2. Data

In this paper we analyzed the species dataset corresponding to the Zosterophylloids *sensu stricto* compiled in Cascales-Miñana and Meyer-Berthaud (2014). To be consistent with the recent contributions of Hao and Xue (2013) and Xiong et al. (2013), minor modifications were included regarding the paleogeographical and stratigraphical distributions of a few taxa (Table S1). The global dataset comprises 28 genera and 93 species.

In accordance with the paleophytogeographical studies carried out thus far on early terrestrial floras, we extracted three sub-datasets corresponding to northeastern Gondwana, Laurussia and Siberia (see Introduction). The northeastern Gondwana dataset includes 39 species from Australia, South China and the Shan–Thai Block. It corresponds to a unit proposed by Hao and Gensel (1998) and supported by several Chinese paleobotanists and vertebrate paleontologists (Pan et al., 1996; Wang et al., 2010; Hao and Xue, 2013). The Laurussian dataset includes 57 species from North America and Europe. Following Edwards and Berry (1991), we analyzed the Siberian dataset (9 species) separately. The Kazakhstan dataset, which is represented by five endemic species from Kazakhstan and Xinjiang, and the northwestern Gondwana dataset, by a single species from Venezuela (*Serrulacaulis* cf. *S. furcatus*; Berry and Edwards, 1994; Xu et al., 2011), have a sample size that is too small for regional data analysis. They were considered only in the analysis of global diversity patterns. See Tables S1–S5 for details.

3. Methods

Several multivariate statistical techniques (cluster analysis and ordination methods) have been used thus far to study the phytogeographic distribution of taxa and delineate phytogeographic units for the Paleozoic (e.g., Raymond et al., 1985; Raymond et al., 2006; Wellman et al., 2013). They were based on datasets representing assemblages on a worldwide basis, with a large number of taxa showing large spatial distributions. Our zosterophylloids dataset comprises too few widely distributed species for using such multivariate statistical methods. Our preliminary attempts using these approaches provided uninterpretable results.

As an alternative, we studied the dynamics of the Zosterophylloids separately for northeastern Gondwana, Laurussia and Siberia, compared them, and estimated how much each of these specific patterns contributed to the global biodiversity dynamics of the group. The methods that we used in this paper are presented below. They are detailed in Cascales-Miñana and Meyer-Berthaud (2014).

3.1. Diversity measures

Analyses of diversity patterns through time were conducted at species and sub-stage levels. Two diversity measures were considered, the total and boundary-crossing standing diversities (Foote, 2000; Alroy et al., 2001). Considering a time interval, the former one corresponds to the total number of species recorded within this interval whereas the latter corresponds to the number of taxa that cross its bottom boundary. The total diversity measure does not provide any information on the ratio of originations versus extinctions at any time interval. Because it includes singletons, i.e., species that originated and became extinct in the same interval, this measure is also sensitive to sampling intensity and to the duration of intervals. In comparison, boundary-crossing standing diversity minimizes such biases. Moreover, this measure indicates changes in the level of originations compared to extinctions between successive time intervals (see Bambach et al., 2004 and references therein). Indeed, comparing the boundary-crossing diversities of two successive time intervals, t_1 and t_2 , is equivalent to estimating the turnover of taxa within t_1 . An increase of boundary-crossing diversity between t_1 and t_2 means that the number of taxa entering t_2 , i. e. leaving t_1 , is greater than that entering t_1 , in other words, that the number of originations exceeds the number of extinctions within t_1 . On the contrary, a decrease of boundary-crossing diversity between t_1 and t_2 means that the number of extinctions exceeds the number of originations within t_1 .

3.2. Sampling intensity measures

Sampling-corrected (so-called residual) diversity curves were used to identify what portions of the total and boundary-crossing

standing diversity curves were not driven by sampling intensity (e.g., Brocklehurst et al., 2013). They were obtained using Lloyd's (2012) method. Lloyd's sampling-driven model assumes that the true taxonomic richness is constant and the observed richness, a consequence of sampling. This method calculates when the taxonomic richness is greater/lower than that expected from sampling. To estimate the sampling effort, we used the minimum number of fossil assemblages (i.e., sensu Raymond and Metz, 1995) recorded for each time interval (Benton et al., 2011). The residual diversity curves were obtained using the R statistical software (R Developmental Core Team, 2013). See <http://www.graemetlloyd.com/methscd.html> for implementation.

4. Results

4.1. Spatial distribution of zosterophylloids

When examining the raw data at species level (Table 1), the most widely distributed taxon, *Margophyton goldschmidtii*, is found in three paleogeographical units, Laurussia, Siberia and northeastern Gondwana; this occurs in the Emsian. During the Pragian, when diversity reaches its highest values in all units, only five well-defined species are shared by two units. Three co-occur in Laurussia and Siberia (*Zosterophyllum rhenanum*, *Sawdonia ornata*, *Distichophytum mucronatum*), one in northeastern Gondwana and Laurussia (*Zosterophyllum australianum*) and one in northeastern Gondwana and Siberia (*Zosterophyllum myretonianum*) (Table 1). From the six zosterophylloid species reported from the Kazakhstan and northwestern Gondwana units, two are endemic (*Sawdonia curstipa*, *Serrulacaulis spineus*), the others being insufficiently preserved to be assigned to a taxon (Table S4).

When examining the raw data at generic level (Table 2) and when excluding *Zosterophyllum*, which is not a natural group, the most widely distributed genera are *Distichophytum* in the Pragian and *Margophyton* in the Emsian; both are found in three paleogeographical units only, Laurussia, Siberia and northeastern Gondwana. During the Pragian, there are three genera shared by Laurussia and northeastern Gondwana (*Hicklingia*, *Distichophytum*, *Gosslingia*) and two by Laurussia and Siberia (*Distichophytum*, *Sawdonia*). It is not before the Givetian that a few taxa (*Sawdonia*, *Serrulacaulis*) co-occur in any of these paleophytogeographic units and Kazakhstan or northwestern Gondwana.

4.2. Global diversity curves

Figs. 2A and 3A show the diversity analyses conducted on the global dataset. The total diversity and boundary-crossing curves show similar trends (Fig. 2A). The highest value of total diversity (44 species) occurs in the late Pragian, that of the boundary crossing diversity (25 species) a little earlier, in the middle Pragian. Both diversity metrics drop down to less than 20 species in the Emsian. Discrepancies between the two

Table 1
Zosterophylloids species shared by at least two phytogeographic units in the Devonian. See supplementary material for details.

Species	Time units			
	Lochkovian	Pragian	Emsian	Eifelian
<i>Margophyton goldschmidtii</i>	Siberia	Siberia	Laurussia Siberia NE Gondwana	–
<i>Zosterophyllum myretonianum</i>	Laurussia	Siberia NE Gondwana	Laurussia	–
<i>Zosterophyllum rhenanum</i>	–	Laurussia Siberia	–	–
<i>Zosterophyllum australianum</i>	–	Laurussia NE Gondwana	NE Gondwana	–
<i>Sawdonia ornata</i>	–	Laurussia Siberia	Laurussia Siberia	Laurussia
<i>Distichophytum mucronatum</i>	–	Laurussia Siberia	Laurussia	–

Table 2

Zosterophylloids genera shared by at least two phytogeographic units in the Devonian. See supplementary material for details.

Genera	Time units				
	Lochkovian	Pragian	Emsian	Eifelian	Givetian
<i>Margophyton</i>	Siberia	Siberia	Laurussia Siberia NE Gondwana	–	–
<i>Hicklingia</i>	–	Laurussia NE Gondwana	–	Laurussia	–
<i>Distichophytum</i>	–	Laurussia Siberia NE Gondwana	Laurussia	–	–
<i>Sawdonia</i>	–	Laurussia Siberia	Laurussia Siberia	Laurussia	Siberia Kazakhstan
<i>Gosslingia</i>	Laurussia NE Gondwana	Laurussia NE Gondwana	–	–	–
<i>Serrulacaulis</i>	–	–	–	Laurussia	Laurussia NW Gondwana Kazakhstan
<i>Oricilla</i>	–	NE Gondwana	Laurussia	–	–

metrics are due to singletons. They are numerous in South China during the late Pragian and in Belgium during the early Emsian.

The boundary-crossing curve shows that originations largely exceeded extinctions in the early Pragian. Extinctions increased in the middle Pragian and became prominent in the late Pragian. A large number of taxa from the latter time interval did not cross the late Pragian–early Emsian boundary. Extinctions also exceeded originations in the late Emsian and early Eifelian.

The residual diversity curve indicates that, at a global scale, the values recorded for the middle Pragian–late Emsian interval are independent from the number of sampled assemblages (Fig. 2A). The middle and late Pragian show significant positive values of species richness. Despite the high number of singletons recorded for this time period, the late Pragian registers the highest diversity value independently from the sampling intensity. Two significant negative peaks are observed in the early and late Emsian. The most important one occurs in the early Emsian confirming the trend shown by the diversity curves (Fig. 2A).

4.3. Laurussian diversity curves

Figs. 2B and 3B show the diversity analyses conducted on the Laurussian sub-dataset. The total diversity and boundary-crossing curves show different trends between the middle Pragian and the middle Emsian (Fig. 2B). For Laurussia, the total diversity curve shows four peaks. The highest one (23 species, with a large contribution of Belgian taxa) occurs in the early Emsian. There are two smaller peaks in the early Pragian (17 species) and the late Emsian (14 species), and a minor peak in the late Eifelian (4 species). The boundary-crossing diversity curve, which excludes the singletons, shows its highest value in the middle Pragian (16 species), and in the middle and late Emsian (13 species). For the boundary-crossing diversity measure, the early Emsian corresponds to a time of low diversity (6 species) rather than a peak as observed for the total diversity.

The trends in the relative levels of originations versus extinctions shown by the boundary-crossing curve are similar to those of the global dataset.

The residual diversity curve shows three “windows” where diversity is independent of the sampling intensity (Fig. 3B). The late Pragian corresponds to a major positive peak whereas the early to middle Lochkovian and the middle Emsian to early Eifelian time intervals show low values of residual diversity. The late Emsian corresponds to the most negative excursion. Residual diversity substantially increases in the early Eifelian, after which no significant change is visible.

The combined results of the total, boundary-crossing and residual diversity curves (Fig. 2B shaded areas, Fig. 3B) indicate that, in Laurussia, zosterophylloids diversity was significantly low in the early and

middle Lochkovian, and high in the late Pragian. A large number of species, however, became extinct in the latter time interval and did not cross the late Pragian–early Emsian boundary. Diversity decreased in the Emsian, the lowest significant value reached in the late Emsian.

4.4. Northeastern Gondwanan diversity curves

Figs. 2C and 3C show the diversity analyses conducted on the northeastern Gondwanan sub-dataset. Compared to the other datasets, the last record of zosterophylloids in northeastern Gondwana occurs early, in the early Eifelian. Despite a small difference in the early Pragian, the total diversity and boundary-crossing curves show comparable trends (Fig. 2C). The highest value of total diversity (25 species) occurs in the late Pragian and is related to the large number of singletons (13 species) documented for the Posongchong assemblages of South China. The highest value of boundary crossing diversity is much lower (8 species) and corresponds to the middle and late Pragian. Diversity decreases to 4 species in the middle and late Emsian.

The boundary-crossing curve shows that originations exceeded extinctions in the early and middle Lochkovian and in the early Pragian. Extinctions were relatively important in the late Lochkovian and the late Pragian. A relatively large number of taxa from the latter time interval did not cross the late Pragian–early Emsian boundary.

The residual diversity curve shows three “windows” where diversity is independent from the sampling intensity (Fig. 3C). A significant rise of diversity occurs between the late Lochkovian and the middle Pragian when the maximum residual value is observed. Statistically significant low values of species richness are observed during the late Pridoli and the middle and late Emsian. Interestingly, the variation of residual diversity remains within confidence limits during the late Pragian confirming that the highest value of total diversity recorded for this dataset in this time interval is artefactual and related to sampling.

The combined results of the total, boundary-crossing and residual diversity curves (Fig. 2C shaded areas, Fig. 3C) indicate that, in northeastern Gondwana, zosterophylloids diversity was significantly low in the late Pridoli and high from the late Lochkovian to the middle Pragian where it reached its maximum value. These results also acknowledge low diversity values in the middle and late Emsian. The highest significant level of diversity registered in northeastern Gondwana during the middle Pragian was moderate compared to that of Laurussia, in the late Pragian.

4.5. Siberian diversity curves

Figs. 2D and 3D show the diversity analyses conducted on the Siberian sub-dataset. The total diversity and boundary-crossing curves show

Fig. 2. Phytogeographic diversity patterns of Zosterophylloids. Comparative analysis of boundary-crossing (bold line) and total (dashed line) diversity curves. Temporal ranges are based on first-last appearance data. (A) Global dataset. (B) Laurussian phytogeographic unit. (C) Northeastern Gondwanan phytogeographic unit. (D) Siberian phytogeographic unit. Data analyses are conducted at species and sub-stage levels. Gray areas indicate the time intervals when diversity is independent from the number of sampled localities (see Fig. 3).

similar trends (Fig. 2D). The total diversity curve shows its maximum value (8 species) in the late Pragian, the boundary-crossing curve, in the middle and late Pragian (6 species). Both diversity curves exhibit a marked decrease of diversity in the early to middle Emsian time interval.

The boundary-crossing curve shows that originations exceeded extinctions in the early Pragian whereas a relatively large number of taxa from the late Pragian did not cross the late Pragian–early Emsian boundary.

The residual diversity curve shows only two “windows” of limited extent where diversity is independent from the sampling intensity (Fig. 3D). A significant positive peak is registered in the middle Pragian and a significant negative peak in the early Emsian. This evidence suggests that the main patterns of diversity change illustrated by the diversity curves are not biased by sampling (Fig. 2D, shaded areas).

5. Discussion

Paleophytogeographical studies are constrained by the intrinsic biases and limitations of the plant fossil record (e.g., incorrect identifications and assessments of systematic relationships, inaccurate dating of non-marine fossil localities, unequal sampling of all stages),

by ambiguities concerning the original habitats and by uncertainties about the geographic configuration of past times (Edwards, 1990; Wnuk, 1996; Wellman et al., 2013). In this paper, we attempted to minimize sampling biases by using a set of methods successfully applied in other diversity analyses (e.g., Foote, 2000; Alroy et al., 2001; Lloyd, 2012). Another possible improvement to increase the robustness of our results would include a complete taxonomic and systematic revision of the group, a long-term task that is largely beyond the scope of this work. The taxic diversity patterns described in this paper therefore refer to apparent spatial diversity changes (diversity fluctuations calculated from fossil data taken at face value, Foote, 2001).

Our results showed that the number of species recorded prior to the Lochkovian for Laurussia or prior to the Pridoli for northeastern Gondwana, and posterior to the early Eifelian, is strongly dependent on the number of sampled localities. The Late Silurian record of zosterophylls, therefore, is unreliable for assessing the early diversity patterns of the group. The Ludfordian genera reported from Bathurst Island (Table S1) may not give a correct indication of when and where the first zosterophylls originated.

The following discussion focusses on a time window that minimizes the biases related to the sampling effort. It compares the patterns

Fig. 3. Model of residual estimates for taxonomic richness of Zosterophyllopsida using the minimum number of sampled localities as sampling proxy. Data analyses conducted at species and sub-stage levels. Dashed lines represent the 95% confidence interval based on the standard deviation of the model. Zero represents the model estimate. Values above zero represent greater than expected richness and below zero lower than expected. Statistical significance is assumed when the residual diversity estimates exceed the upper and/or lower boundaries of the confidence interval of proxy-biased model (Lloyd, 2012). (A) Global dataset. (B) Laurussian phytogeographic unit. (C) Northeastern Gondwanan phytogeographic unit. (D) Siberian phytogeographic unit. Significant time intervals are indicated by arrows.

of zosterophyllopsid diversity of Laurussia, Siberia, northeastern Gondwana, northwestern Gondwana and Kazakhstan during the Early Devonian and examines what factors may explain their similarities and differences.

5.1. Compared patterns of diversity changes between phytogeographic units

The Zosterophyllopsida of Laurussia, northeastern Gondwana and Siberia show similar curves of diversity changes (Fig. 2B–D). Their specific diversity increased markedly during the Lochkovian. This phase of radiation reached a peak that, depending on the phytogeographic unit and the approach taken to measure diversity, occurred between the middle Pragian and the early Emsian (Table 3). The peak was followed by an abrupt drop of diversity ending in the Eifelian (Fig. 2; see also Cascales-Miñana and Meyer-Berthaud, 2014). For all regions, the boundary diversity curves provide evidence of a high level of

extinction in the late Pragian. This similarity in the shape of the diversity curves of the three units suggests that the processes governing these changes, especially the onset of a phase of high extinction just after the phase of radiation, were similar and probably global.

As zosterophyllopsid specific diversity started to decrease, lycopsids and basal euphyllophytes showed the reverse tendency. These trends, quantified in Niklas et al.'s (1985) analysis of land plant diversification, have been confirmed since then with the description, in Laurussia and South China especially, of many new taxa of Pragian and Emsian age belonging to these latter groups (Hao and Xue, 2013, Table A4; Taylor et al., 2009). Disparity also rose significantly in the lycopsids and the basal euphyllophytes. Leaf-like organs diversified in both groups as early as the Pragian (microphylls in the lycopsids, webbed lateral branched systems in the euphyllophytes; Hao and Beck, 1993; Gerrienne et al., 2014), and a primitive type of wood evolved in the euphyllophytes (Gerrienne et al., 2011). In a context of biotic competition, new developmental innovations leading to the evolution, in the

Table 3
Time intervals corresponding to the highest and lowest diversity values of Zosterophyllopsida in Laurussia, northeastern Gondwana, Siberia and for the global dataset, based on total, boundary-crossing and sampling corrected diversity estimates. See supplementary material for details.

	Diversity estimates					
	Highest diversity values			Lowest diversity values following the positive peaks		
	Total (singletons and sampling effect)	Boundary crossing (no singletons)	Sampling corrected	Total (singletons and sampling effect)	Boundary crossing (no singletons)	Sampling corrected
Laurussia	Early Emsian	Middle Pragian	Late Pragian	Middle Eifelian	Middle Eifelian	Late Emsian
Northeastern Gondwana	Late Pragian	Middle and late Pragian	Middle Pragian	Early Eifelian	Early Eifelian	Middle and late Emsian
Siberia	Late Pragian	Middle and late Pragian	Middle Pragian	Early Eifelian	Early Eifelian	Early Emsian
Global	Late Pragian	Middle Pragian	Late Pragian	Early Eifelian	Early Eifelian	Early Emsian

Fig. 4. Comparison between the total diversity curve of Zosterophylloids, the Haq and Shutter (2008, Fig. 2) sea-level curve, and the Joachimski et al.'s (2009, Fig. 7) paleotemperature curve during the Lochkovian-basal Eifelian time interval. Temporal scale adapted from Gradstein et al. (2004).

lycopsids and the euphyllophytes, of specialized taxa with increased efficiency in water transport and photosynthesis may have been disadvantageous for the zosterophylloids.

Climate trends for the Devonian have been reconstructed from several approaches including calculations from the oxygen isotopic composition of brachiopod shells and conodonts (Joachimski et al., 2009) and computer simulation using geochemical models (Simon et al., 2007; Le Hir et al., 2011; Godderis et al., 2014). Joachimski et al.'s (2009) paleotemperature curve calculated at a time resolution roughly comparable to that of our analyses indicates that mean sea water temperatures in tropical to subtropical paleolatitudes, where most zosterophylloids lived, showed a globally steady decrease from the Lochkovian to the Eifelian (Fig. 4). It is possible that from the late Pragian onwards, when the level of extinctions started to be significant, the decreasing temperatures were no longer optimal for the Zosterophylloids.

Sea-level was low throughout the Early Devonian. The short-term curve reconstructed by Haq and Shutter (2008), however, shows that successive sea-level changes of high magnitude occurred during the Pragian, a time interval characterized by its short duration (about 4.5 million years according to the stratigraphic chart used by Haq and Shutter; about 3.2 million years according to the 2014 chart produced by ICS — <http://stratigraphy.org/index.php/ics-chart-timescale>) (Fig. 4). If, as hypothesized by Hotton et al. (2001) the Zosterophylloids preferred stable habitats, the magnitude and frequency of the water table changes in the Pragian may have led to disturbances in both coastal environments and inland that were detrimental to the group and resulted in the first elevated ratio of extinctions versus originations at the end of this time interval.

We admit that hypotheses involving external factors such as temperature and sea-level changes to explain the highest level of extinctions at the end of the Pragian are highly conjectural, hardly testable, and not entirely satisfactory. Invoking the marked sea-level changes and the resulting disturbances that occurred throughout the Pragian as causing the extinctions of the group at the end of the interval contradicts the peak of diversity observed just before in the same time-interval. Moreover, a clade-based partitioning of early Devonian landscapes has been challenged by Kennedy et al.'s (2012) recent analysis of the distribution of plants in an Emsian locality of New Brunswick. These authors show that zosterophylls occur in the same beds as

lycopsids and euphyllophytes with little evidence for ecological preferences in terms of stable versus disturbed habitats, and wetlands versus drier sites, for any of these groups. Assessing causal relationships between the diversity patterns of the Zosterophylloids and climatic or eustatic factors requires a deeper understanding of the ecological requirements of these plants together with a finer resolution of the spatial and temporal changes of these external factors.

If the patterns of diversity change are broadly comparable for the Laurussian, northeastern Gondwanan and Siberian paleophytogeographic units, the time occurrences of their respective highs and lows, however, differ (Table 3). With the sampling corrected approach, the highest diversity value for Laurussia occurs in the late Pragian, that for northeastern Gondwana and Siberia a little earlier, in the middle Pragian. The same difference is observed with the total diversity approach, which reveals a peak in the early Emsian for Laurussia and in the late Pragian for northeastern Gondwana and Siberia; it is less marked with the boundary crossing approach where the highest positive peak of diversity occurs in the middle Pragian for the three subunits. Following the radiation, the lowest values of diversity recorded with the sampling corrected approach are found in Siberia first (early Emsian) then northeastern Gondwana (middle Emsian) and finally Laurussia (late Emsian). A similar difference between Laurussia and the two other paleophytogeographic units is observed again when the two other approaches are used. The highest peaks of diversity for the global dataset occur in the same substages as those for Laurussia, a result that may have been anticipated when considering the relative sizes of the Laurussian, northeastern Gondwanan and Siberian subdatasets. The lowest diversity values, however, would have been more difficult to predict. Indeed, the early Emsian low value indicated by the sampling corrected approach appears relatively early and predicts a dramatic drop of diversity when considering that the same approach identifies the immediate previous substage (late Pragian) as the interval showing the highest diversity value. The early Eifelian values registered with the total and boundary crossing approaches correspond to a less spectacular decrease of diversity.

To summarize, the radiation of the Zosterophylloids that started in the Lochkovian reached a peak in the Pragian, earlier in northeastern Gondwana and Siberia than Laurussia. Diversity declined markedly in the Emsian to reach its lowest values before extinction in northeastern Gondwana and Siberia first. These differences provide evidence that the common pattern of diversity changes outlined for all zosterophylloid subdatasets showed some adjustments in relation to paleogeography. The radiation of the Zosterophylloids ceased earlier in northeastern Gondwana (China essentially) or Siberia than Laurussia for reasons that remain to be determined. We argued above that competitive replacement by lycopsids and basal euphyllophytes may have triggered the decrease in diversity of the zosterophylls. In this context, it is interesting to note that the earliest basal euphyllophytes with either planate or laminate lateral branched systems resembling proto-leaves were found in South China (Hao and Beck, 1993; Gerrienne et al., 2014).

5.2. Endemic distribution of the Zosterophylloids in the Early Devonian

The five paleophytogeographic units examined in this paper are characterized by a high number of endemic zosterophyll taxa. When the total observed diversity reaches its highest values, in the Pragian–Emsian time interval, the most widely distributed species (*M. goldschmidtii*) and genera (*Margophyton*, *Distichophyton*) are few and they are shared by three units only, Laurussia, Siberia and northeastern Gondwana (Tables 1–2).

At species level and when excluding taxa referred to as “cf.” or “aff.”, Laurussia and Siberia are the regions that have the highest number of species in common (Table 1). In the Pragian there are three such species (*S. ornata*, *D. mucronatum* and *Z. rhenanum*). At generic level, Laurussia shares more genera with northeastern Gondwana (i. e.

Hicklingia, *Distichophytum* and *Gosslingia* in the Pragian) than with Siberia (*Distichophytum* and *Sawdonia* in the Pragian) (Table 2). Whatever the taxic level examined, however, the number of shared taxa is too small to make any inference about preferential levels of exchanges and proximity between paleogeographic regions.

The pronounced endemism of the Zosterophylloids supports the broad patterns of phytogeographical differentiation hypothesized by previous authors for the Late Silurian and the Early Devonian (Petrosyan, 1967; Raymond, 1987; Edwards and Berry, 1991; Raymond et al., 2006; Wellman et al., 2013). It indicates that the floral exchanges that resulted in the settlement of common taxa between Laurussia, Siberia, northeastern Gondwana, Kazakhstan and northwestern Gondwana were very limited and that geographic barriers may have prevented massive gene flow between these regions. This result is consistent with biogeographical analyses conducted on macrovertebrates (Young, 2010). Differences in the distribution patterns of fishes in the Late Silurian and the Early Devonian have led to the recognition of several distinctive faunal provinces. Laurussia, Siberia, Kazakhstan and Gondwana were interpreted as landmasses separated by prominent seaways that became narrower in the Middle and Late Devonian (Young, 2003; Zhao and Zhu, 2010). These studies also distinguish a Pan-Cathaysian province documenting the connection between south China, north China and Tarim (Young and Janvier, 1999), but suggesting the separation of China and Australia in the Late Silurian and the Early Devonian. Because of an insufficient number of plant records from Australia, this view, which contradicts the prevailing hypotheses in paleobotany (Hao and Gensel, 1998; Hao and Xue, 2013), was not tested in this paper.

6. Conclusion

The Silurian record of the Zosterophylloids is biased and significantly dependent on the number of sampled localities. The actual time and location of origination of the Zosterophylloids is uncertain.

If the available zosterophylloid dataset is too small, in the number of taxa or localities sampled, for identifying paleophytogeographic units, it allows acknowledgement of the taxic distinctiveness of each of the paleogeographic units hypothesized for the Early Devonian (Petrosyan, 1967; Raymond, 1987; Edwards and Berry, 1991). It recognizes that the zosterophylloid species of Laurussia, Siberia, northeastern Gondwana, Kazakhstan and northwestern Gondwana show little overlap. This confirms the high level of endemism shown by other organisms such as fishes in the Early Devonian (Young, 2003; Zhao and Zhu, 2010) and provides a new line of evidence for substantial marine barriers separating the corresponding landmasses at this time.

Analysis of the global dataset shows that, after a period of low diversity spanning the Late Silurian–early Lochkovian, the Zosterophylloids underwent a radiation that culminated in the Pragian then declined from the early Emsian until the Late Devonian. Despite their taxic differences, the specific diversities of the Zosterophylloids in Laurussia, northeastern Gondwana and Siberia show comparable curves, all characterized by a similar phase of origination, a high level of extinction in the late Pragian, and a marked drop of diversity thereafter. These similarities suggest that the changes may have been driven by the same factors in the three regions.

We hypothesize that the high level of extinctions recorded in the late Pragian may have resulted from competition with the rising lycopsids and basal euphyllophytes. These groups evolved new characters such as wood and diversified leaf-like organs that led to an increase of specialized taxa with enhanced performances in fundamental functions such as photosynthesis. The decline of the zosterophylloids may have been precipitated by a global context of environmental changes.

Radiation of the Zosterophylloids stopped earlier in northeastern Gondwana (China essentially) and Siberia than Laurussia. If our hypothesis about the competitive replacement of the Zosterophylloids is

confirmed, we anticipate an earlier burst of the lycopsids and basal euphyllophytes in the two former regions than in Laurussia.

Supplementary data to this article can be found online at <http://dx.doi.org/10.1016/j.palaeo.2015.01.024>.

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